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## Abstract

Transmission spectroscopy in the primary transit of an exoplanet has proven to be very useful for obtaining information of exoplanet atmospheres from both ground-based facilities and space telescopes. The Calar Alto high-Resolution search for M dwarfs with Exoearths with Near-infrared and optical Échelle Spectrographs (CARMENES) instrument has started being operative in early 2016 and here, we explore its capabilities for extracting information about Hot Jupiter atmospheres taking advantage of its ultra-stability, wide spectral interval (0.52-1.7  $\mu\text{m}$ ), and high spectral resolution ( $R=82000$ ). We present some preliminary results of our simulations of the primary transit transmission spectra of HD 189733b in the 1-1.7  $\mu\text{m}$  spectral range where several molecules, such as water vapour, carbon monoxide, carbon dioxide and methane, have strong ro-vibrational bands. Sensitivity studies are presented for the range of expected concentrations of these species, as well as for the expected range of temperature profiles. Our simulations have been performed using the line-by-line Karlsruhe Optimized and Precise Radiative Transfer Algorithm (KOPRA) adapted for exo-atmospheres.

## 1. Introduction

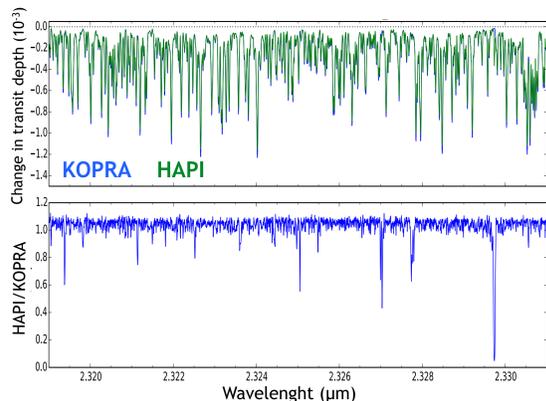
- HD 189733b is a Hot Jupiter exoplanet orbiting an orange dwarf star (K1.5V). It is one of the most studied exoplanets and, in the next months, the CARMENES instrument will obtain high resolution primary transit spectra. These will be used to search for atmospheric molecules such as H<sub>2</sub>O, CO and CH<sub>4</sub>.
- Here we describe the tools and methods we use for the simulation of its transmission spectra for several P,T profiles and molecule volume mixing ratios ('vmr': abundance of one component of a mixture relative to that of all other components) and present results on: a) validations of our radiative transfer calculations; b) show the effects of changing p-T profiles and vmr; c) simulations of low-resolution HST/WFC3 observations and d) simulations of high-resolution CARMENES data in the same spectral range.

Table 1.2 CARMENES Parameters	
Wavelength coverage ( $\Delta\lambda$ )	VIS: 0.52-0.96 $\mu\text{m}$ , NIR: 0.96-1.7 $\mu\text{m}$
Spectral Resolution (R)	VIS: 94,600; NIR: 80,400
Working Temperature ( $T_{\text{work}}$ )	VIS: 285.00 $\pm$ 0.05 K; NIR: 140.00 $\pm$ 0.05 K

## 3. Comparison between KOPRA and a HAPI-based code calculations

- Hitran Application Programming Interface (HAPI): (a) Line-by-line data from HITRANonline; (b) computes high-resolution spectral simulations (absorption coefficients and radiance, transmission and absorption spectra); (c) Choice of Instrumental Line Shape (ILS) to simulate experimental data.

Fig3: Comparison between both codes for a water vapor simulation in an isothermal atmosphere 1500K with a vmr of  $1 \cdot 10^{-4}$  and a boxcar ILS (0.01  $\text{cm}^{-1}$ ).



- Good agreement between KOPRA and HAPI calculations. Both codes calculate almost the same spectral transit depth.

- Differences:
  1. Absorption base level: Due to different atmospheric grids and  $R_p$  values used.
  2. Missing line in HAPI code (2.329 $\mu\text{m}$ ).
  3. Small differences (<5%) in transit depth of weak absorption lines.

## 5. Modeling the HST/WFC3's HD 189733b Transmission Spectra

- Theoretical atmospheric models used to simulate the HST/WFC3 (FWHM=100  $\text{cm}^{-1}$ ) data (based on [4]).

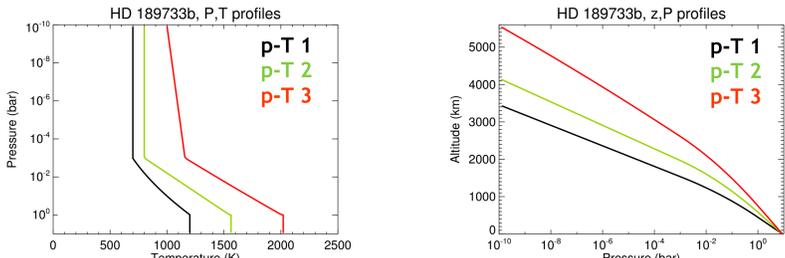


Fig5.1: Atmospheric p-T profiles used in the simulations in Figs 5.2, 6.1 and 6.2. Profiles 2 and 3 similar to dayside temperature profiles in [4]. Profile 1 is colder, more suited for the HST/WFC3 transmission measurements (see below).

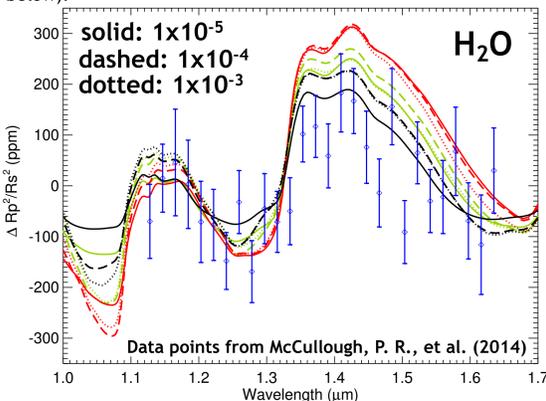


Fig5.2: Water vapor transmission spectra of HD 189733b as measured by HST/WFC3 compared to our theoretical models. Solid, dashed and dotted lines are for vmr's of  $1 \cdot 10^{-5}$ ,  $1 \cdot 10^{-4}$ ,  $1 \cdot 10^{-3}$ , respectively. Colors indicate the temperature profile as in Fig.5.1.

- Dayside P,T profiles of [4] are too hot to match the HST/WFC3 transmission data for the planetary limb atmosphere. Maybe this is partly because of the use of different spectroscopic linelists.

- The best fits to the data are given by the colder profile and vmr of  $1 \cdot 10^{-5}$  (solid black line) or vmr of  $1 \cdot 10^{-4}$  (dashed black line). A vmr of  $6.3 \cdot 10^{-6}$  as suggested by [6] might be too low to reproduce the peak-to-valley amplitude (perhaps due to different linelists).

- Large cross-talk between the p-T profile and the H<sub>2</sub>O abundance (see Fig.6.2).

## 7. Conclusions and Future Work

- Reduction of the the peak-to-valley amplitude using HITEMP 2010 compared to HITRAN 2012.
- Dayside T,P profiles from [4] are too hot to match the HST/WFC3 observations of the limb atmosphere.
- The best fits to the HST/WFC3 data are given by the colder p-T profile and H<sub>2</sub>O abundance of  $1 \cdot 10^{-5}$  or  $1 \cdot 10^{-4}$ .
- The large cross-talk between the p-T profile and H<sub>2</sub>O abundance is reproduced in our HST/WFC3 simulations.
- Molecular detection of CO in the high-resolution transmission spectra of CARMENES might be possible.
- CARMENES might help decrease the large cross-talk between T and water vapour abundance.

- Study different P,T profiles to find the best fits to HST/WFC3 data.
- Characterize the telluric contamination of the Earth's atmosphere and its removal.
- Perform similar studies for other exoplanets, such as HD 209458b.

## 8. References

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- [4] Madhusudhan, N., & Seager, S. (2009). *The Astrophysical Journal*, 707(1), 24.
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- [6] Madhusudhan, N., et al. (2014). *The Astrophysical Journal Letters*, 791(1), L9.

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## 2. Tools and Methods

### 2.1 Concept of 'Change in Transit Depth' ([1] and [2]):

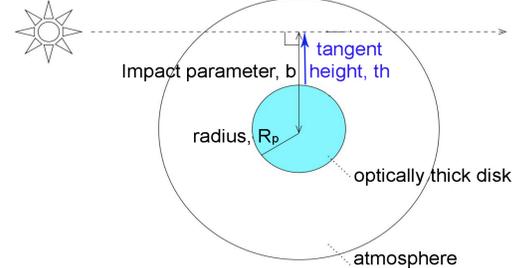


Fig.2 adapted from [1]: Transmission of starlight through the planetary limb. The planet itself (in blue,  $p > 1$  bar) is optically thick at all wavelengths. The scale is distorted for clarity.

- See [1] for a complete derivation of the magnitude Change in transit depth

$$ctd = - \frac{\sum_{\lambda} \pi R_p^2}{\pi R_s^2}, \quad \text{Eq. 1}$$

where the atmospheric absorption is:

$$\sum_{\lambda} = \int_{R_p}^{b_{\text{max}}} 2\pi b db [1 - e^{-\tau_{\lambda}(b)}] \quad \text{Eq. 2}$$

and  $b = R_p + \text{tangent height ('th' above the "solid surface", see Fig.2)}$ .

### 2.2 Karlsruhe Optimized and Precise Radiative Transfer Algorithm, KOPRA [3]:

- Line-by-line radiative transfer model
- Kuntz implementation of the Humlicek algorithm for the Voigt line-shape.
- Interface for generic NLTE-model GRANADA (switched off for this work)
- Supports vibrational and rotational non-LTE
- Computes spectra and Jacobians for LTE and non-LTE
- Accuracy of absorption coefficient calculation (and number of lines)
- Accuracy of frequency integration
- Interfering species to be included (and altitude range)

## 4. Comparison between spectroscopic databases

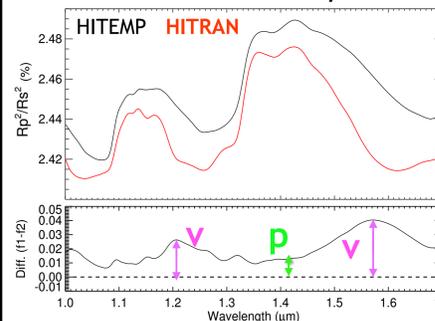


Fig4: Comparison of the spectrum ratio (-Eq.1 +  $R_p^2/R_s^2$ , as described in [1]) obtained using the HITEMP 2010 and HITRAN 2012 spectroscopic databases. An isothermal atmosphere at 1500K, H<sub>2</sub>O vmr of  $1 \cdot 10^{-4}$  and a gaussian ILS of FWHM=100  $\text{cm}^{-1}$  have been used.

- HITEMP includes more lines and therefore is more absorbent.
- Note that the peak to valley amplitude is significantly reduced using HITEMP due to stronger absorption in the valleys (pink arrow) with respect to the band heads (green arrow).

## 6. Simulation of the transmission as seen by CARMENES

- Model atmospheres in Fig.5.1 are used to simulate the transmission spectrum with the CARMENES high spectral resolution (see Table 1.2) in order to explore its capability to a) identify trace gases as CO, and b) constrain the p-T and H<sub>2</sub>O abundance.

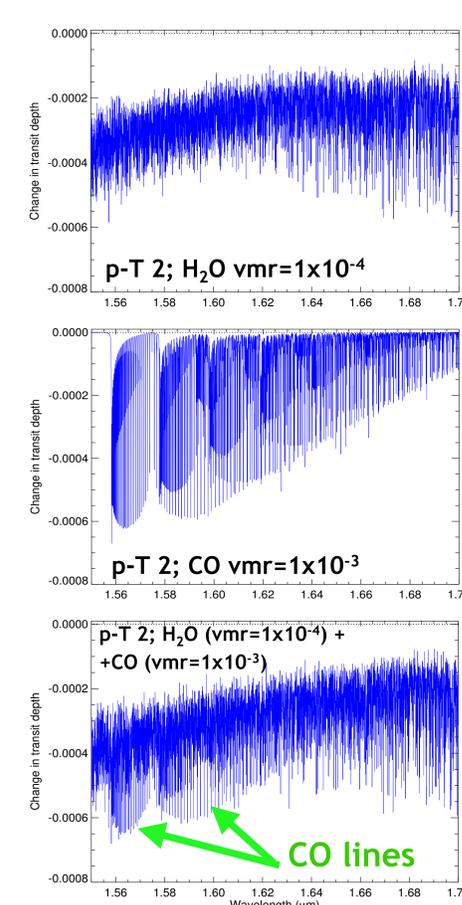


Fig6.1: Simulated transmission spectra of HD 189733b at the CARMENES spectral resolution of 0.1  $\text{cm}^{-1}$  for the intermediate p-T profile 2, H<sub>2</sub>O vmr of  $1 \cdot 10^{-4}$  and CO vmr of  $1 \cdot 10^{-3}$ .

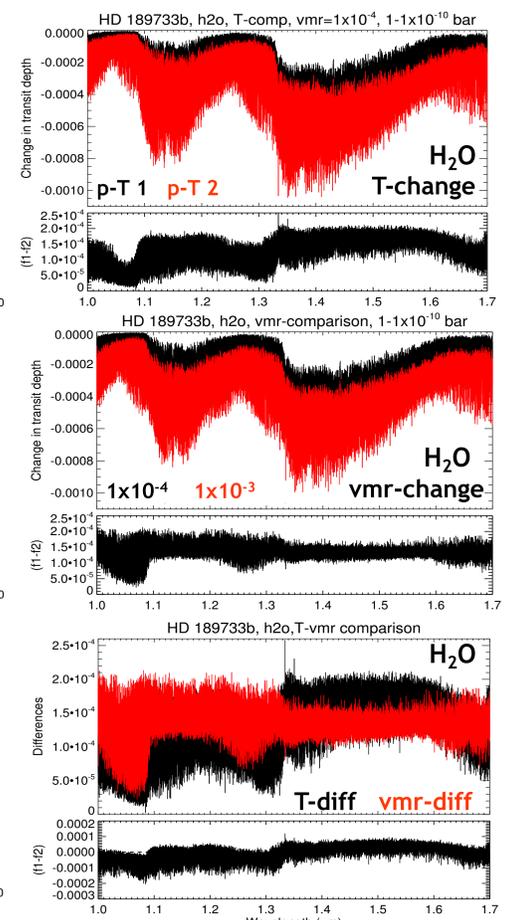


Fig6.2: Analysis of the effect of changing temperature or H<sub>2</sub>O volume mixing ratio.

- Changing vmr increases the absorption uniformly. Changing T has a stronger impact on the deepest absorption band. This could help decrease the huge cross-talk between T (several hundred K) and H<sub>2</sub>O abundance (spanning 3 orders of magnitude).

- The CARMENES high-resolution transmission spectrum likely allows us to identify CO lines along with water vapour.